

SHORTER CONTRIBUTIONS

THE DISTRIBUTION OF RAINDROPS WITH SIZE

By J. S. Marshall and W. Max. Palmer¹
 McGill University, Montreal
 (Manuscript received 26 January 1948)

Measurements of raindrop records on dyed filter papers were made for correlation with radar echoes (Marshall, Langille, and Palmer, 1947) These measurements have been analyzed to give the distribution at drops with size (fig. 1). The distributions are in fair agreement with those of Laws and Parsons (1943).

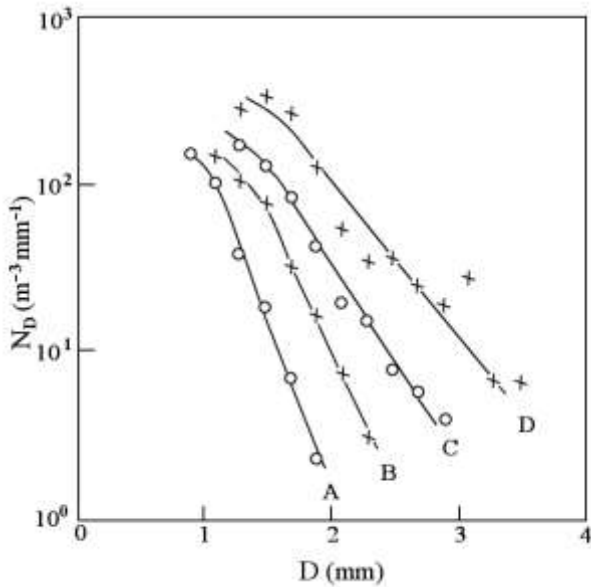


FIG. 1. Distribution of number versus diameter for raindrops recorded at Ottawa, summer 1946. Curve A is for rate of rainfall 1.0 mm hr⁻¹, curves B, C, D, for 2.8, 6.3, 23.0 mm hr⁻¹. $N_D \delta D$ is the number of drops per cubic meter, of diameter between D and $D + \delta D$ mm. Multiplication by 10^{-6} will convert N_D to the units of equation (2).

Except at small diameters, both sets of experimental observations can be fitted (fig. 2) by a general relation,

$$N_D = N_0 e^{-\Lambda D} \quad (1)$$

where D is the diameter, $N_D \delta D$ is the number of drops of diameter between D and $D + \delta D$ in unit volume of space, and N_0 is the value of N_D for $D = 0$.

It is found that

$$N_0 = 0.08 \text{ cm}^{-1} \quad (2)$$

for any intensity of rainfall, and that

$$\Lambda = 41R^{-0.21} \text{ cm}^{-1}, \quad (1)$$

where R is the rate of rainfall in mm hr⁻¹.

For diameters less than about 1.5 mm, both sets of observations fall short of the value for N_D given by equation (1), and they disagree slightly with each other. Laws and Parsons' observations are better in

this region, and tend toward a common value of N_0 for all rates of rainfall.

The mass of rain water M per unit volume of space, and the sum Z of sixth powers of drop diameters in unit volume (a radar quantity), can be calculated as functions of Λ from equation (1), and so correlated with the rate of rainfall R by equation (3). It is of interest to compare these correlations with those obtained when M , Z , and R are determined more directly from the experimental records (table 1). The deficit of

TABLE 1. $M = \frac{1}{6} \pi \Sigma N_D D^3 \delta D$ and $Z = N_D D^6 \delta D$ as functions of the rate of rainfall R

Reference	M mgm m ⁻³	Z mm ⁶ m ⁻³
Marshall, Langille and Palmer (1947)	80 $R^{0.85}$	190 $R^{1.72}$
Revision of the above	72 $R^{0.88}$	220 $R^{1.60}$
Z/R correlation by Wexler (1947) (data of Laws and Parsons (1943))	68 $R^{0.88}$	320 $R^{1.44}$
From equations (1) and (3)	89 $R^{0.84}$	296 $R^{1.47}$

small drops in the observations, as compared with equation (1), should make the observed value of M , and to a lesser extent that of Z , smaller than those derived from the equations.

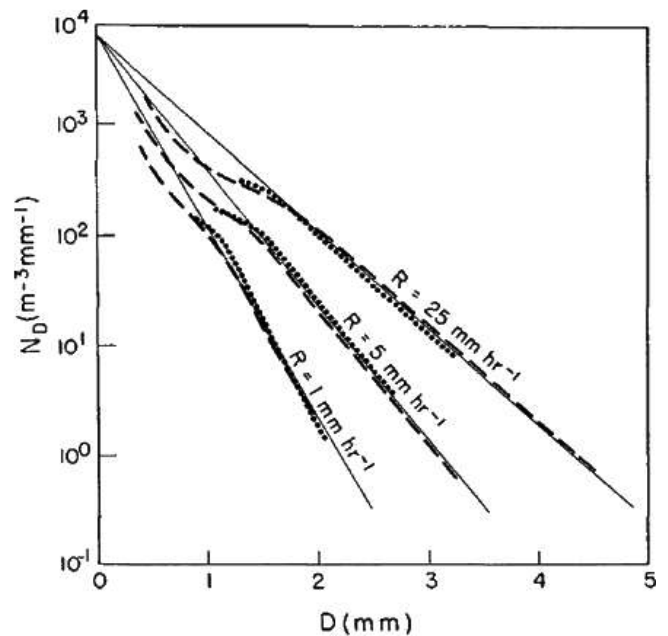


FIG. 2. Distribution function (solid straight lines) compared with results of Laws and Parsons (broken lines) and Ottawa observations (dotted lines).

¹ Holding a Bursary of the National Research Council of Canada.

The exponential distribution of equation (1) is the type that would obtain if growing drops were in continual danger of disintegrating, the likelihood of disintegration being proportional to the increment in diameter or in distance of fall through cloud. Such behavior might be explained by the random accumulation by each drop of electrical charge as more and more randomly charged cloud drops or smaller raindrops are acquired by coalescence, and the resultant disintegration of overcharged drops. Relevant calculations and experiments on coalescence are in progress.

Part of the work reported here was done during summer employment in the Radar Meteorology Section of the Defense Research Board's Radio Propagation Laboratory at Ottawa.

REFERENCES

- Laws, J. O., and D. A. Parsons, 1943: The relation of raindrop-size to intensity. *Trans. Am. geophys. Union*, 24, part II, 452-460.
- Marshall, J. S., R. C. Langille, and W. McK. Palmer, 1947: Measurement of rainfall by radar. *J. Meteor.*, 4, 186-192.
- Werler, R., 1947: Radar detection of a frontal storm 18 June 1946. *J. Meteor.*, 4, SM.

